

## Drought risk in the North Eurasian regions: assessment of El-Nino effects

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We analyzed the probability of drought conditions in the North Eurasian regions and assessed El-Nino/Southern Oscillation (ENSO) effects [1]. Based on observations [2] for the period 1891-2013, we analyzed the spring-summer (May-July) anomalies of surface air temperature (SAT)  $\delta T$  and precipitation  $\delta P$ , as well as of drought index  $D$  in the mid-latitudes of the European (ER) and Asian (AR) parts of Russia .

To estimate the El-Nino/La-Nina effects, we used their indices based on sea surface temperature (SST) in the Niño-3 (150°–90°W, 4°N–4°S), Niño-3,4 (170°–120°W, 4°N–4°S) and Niño-4 (160°E–150°W, 4°N–4°S) regions in the equatorial latitudes of the Pacific Ocean ([ftp://www.coaps.fsu.edu/pub/JMA\\_SST\\_Index/](ftp://www.coaps.fsu.edu/pub/JMA_SST_Index/)). The El-Nino ( $E$ ) and La-Nina ( $L$ ) phases were distinguished using 5-month moving averages of the SST anomaly in the Niño-3 region (JMA index). El-Nino (warm) and La-Nina (cold) phases were defined by the index values of at least 0.5°C and at most –0.5°C, respectively, over six consecutive months (including October–December). All other cases were considered as neutral phases ( $N$ ).

The beginning of 2015 in the El-Nino phase was characterized by the highest positive anomalies in the Niño-4 region and we present here the results with Niño-4 as the El-Nino index. According to ensemble model forecasts, the probability that El-Nino phase will persist till the end of this year is about 60% (more than 70% for May-July and June-August). The corresponding probability for neutral and La-Nina phases is estimated at about 30% and 10%, correspondingly (<http://iri.columbia.edu>).

Table 1 shows the probabilities of positive spring–summer temperature anomalies  $\delta T$  in the ER and AR for different El-Nino phase transitions estimated with the use of the Niño-4 index. The transitions characterized by the highest probability of temperature anomalies, including extreme ones ( $\delta T > 1^\circ\text{C}$ ), for the ER and AR are shown in bold. It should be noted that the  $E \rightarrow E$  transition expected in 2015 is characterized by the highest probability (0.71) of extremely high temperature in May-July over the AR with a low risk of high temperatures for the ER.

**Table 1.** Probability of different surface temperature anomalies ( $\delta T$ ) in the ER (and AR) in May-July for different ENSO phases (characterized by the Niño-4 index) based on observations for 1891–2013

$\delta T$	$N \rightarrow$ $n = 67$			$E \rightarrow$ $n = 28$			$L \rightarrow$ $n = 28$		
	$N \rightarrow E$ $n = 17$	$N \rightarrow L$ $n = 9$	$N \rightarrow N$ $n = 41$	$E \rightarrow E$ $n = 7$	$E \rightarrow L$ $n = 8$	$E \rightarrow N$ $n = 13$	$L \rightarrow E$ $n = 4$	$L \rightarrow L$ $n = 10$	$L \rightarrow N$ $n = 14$
$>0$	0.41 (0.47)	0.56 (0.44)	0.63 (0.61)	0.29 <b>(0.71)</b>	<b>0.88</b> (0.62)	0.54 (0.38)	0.50 (0.50)	0.40 (0.50)	0.43 (0.57)
	0.53 (0.51)			<b>0.57</b> <b>(0.57)</b>			0.44 (0.52)		
$>1 \text{ K}$	0.12 (0.24)	0.22 (0.11)	0.22 (0.22)	0.14 <b>(0.71)</b>	<b>0.50</b> (0.12)	0.31 (0.15)	0.25 (0.25)	0.20 (0.20)	0.21 (0.29)
	0.19 (0.19)			<b>0.32</b> <b>(0.33)</b>			0.22 (0.25)		

Table 2 presents the probability of negative precipitation anomalies in the ER and AR ( $\delta P$ ) in May–July for different ENSO phase transitions estimated with the use of Nino-4 index.

**Table 2.** Probability of negative precipitation anomalies ( $\delta P$ ) in the ER (and AR) in May–July for different ENSO phases (characterized by the Nino-4 index) from observations for 1891–2013

$\delta P$	$N \rightarrow$ $n = 67$			$E \rightarrow$ $n = 28$			$L \rightarrow$ $n = 28$		
	$N \rightarrow E$ $n=17$	$N \rightarrow L$ $n=9$	$N \rightarrow N$ $n=41$	$E \rightarrow E$ $n=7$	$E \rightarrow L$ $n=8$	$E \rightarrow N$ $n=13$	$L \rightarrow E$ $n=4$	$L \rightarrow L$ $n=10$	$L \rightarrow N$ $n=14$
< 0	0.47 (0.35)	0.56 (0.33)	0.51 (0.34)	0.29 (0.57)	0.38 (0.50)	0.54 (0.62)	0.50 ( <b>0.75</b> )	<b>0.60</b> (0.50)	0.43 (0.36)
	<b>0.51</b> (0.34)			0.40 ( <b>0.56</b> )			<b>0.51</b> (0.54)		

Table 3 shows probability of different drought conditions (index  $D$ ) in the ER (and AR) in May–July for different ENSO phase transitions with the use of the Nino-4 index.

**Table 3.** Probability of different drought conditions (index  $D$ ) in the ER (and AR) in May–July for different ENSO phases (characterized by the Nino-4 index) from observations for 1891–2013

$D$	$N \rightarrow$ $n = 67$			$E \rightarrow$ $n = 28$			$L \rightarrow$ $n = 28$		
	$N \rightarrow E$ $n=17$	$N \rightarrow L$ $n=9$	$N \rightarrow N$ $n=41$	$E \rightarrow E$ $n=7$	$E \rightarrow L$ $n=8$	$E \rightarrow N$ $n=13$	$L \rightarrow E$ $n=4$	$L \rightarrow L$ $n=10$	$L \rightarrow N$ $n=14$
$\geq 20\%$	0.24 (0.35)	<b>0.56</b> (0.33)	0.51 (0.29)	0.43 ( <b>0.71</b> )	0.50 (0.38)	0.46 (0.38)	0.25 (0.50)	0.30 (0.30)	0.29 (0.36)
	0.44 (0.32)			<b>0.46</b> ( <b>0.49</b> )			0.28 (0.39)		
$\geq 30\%$	0.12 (0.06)	0.22 (0)	0.27 (0.20)	0 ( <b>0.43</b> )	<b>0.38</b> (0.12)	0.31 (0.23)	0.25 (0.25)	0.30 (0.20)	0.21 (0.21)
	0.20 (0.09)			0.23 ( <b>0.26</b> )			<b>0.25</b> (0.22)		

According to Table 3 the  $E \rightarrow E$  transition expected in 2015 is characterized by the high risk of drought conditions in May–July in AR. Severe drought conditions ( $D \geq 30\%$ ) in AR were realized three times in seven  $E \rightarrow E$  transitions since 1891.

## References

1. Mokhov I.I. and A.V. Timazhev (2013) Climatic anomalies in Eurasia from El-Nino/La-Nina effects. *Doklady Earth Sci.*, **453**(1), 1141-1144.
2. Meshcherskaya A.V., V.M. Mirvis, and M.P. Golod (2011) The drought in 2010 against the background of multiannual changes in aridity in the major grain-producing regions of the European part of Russia. *Tr. MGO*, **563**, 94–121 (in Russian)